

NeQuick model and data assimilation

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> International School of Space Science THE POLAR UPPER ATMOSPHERE: FROM SCIENCE TO OPERATIONAL ISSUES L'Aquila, 16-21 Sep 2018

NeQuick

 The NeQuick is an ionospheric electron density model developed at the T/ICT4D Laboratory of The Abdus Salam International Centre for Theoretical Physics (ICTP), Trieste, Italy, and at the Institute for

Geophysics, Astrophysics and Meteorology (IGAM) of the University of Graz, Austria.

- It is a quick-run empirical model particularly designed for trans-ionospheric propagation applications, conceived to reproduce the median behavior of the ionosphere ("climate").
- <u>http://t-ict4d.ictp.it/nequick2</u>





NeQuick 2

- The model profile formulation includes 6 semi-Epstein layers with modeled thickness parameters and is based on anchor points defined by foE, foF1, foF2 and M(3000)F2 values.
- These values can be modeled (e.g. ITU-R coefficients for foF2, M(3000)F2) or experimentally derived.
- NeQuick inputs are: position, time and solar flux; the output is the electron concentration at the given location and time.
- NeQuick package includes routines to evaluate the electron density along any "ground-to-satellite" ray-path and the corresponding Total Electron Content (TEC) by numerical integration.

B. Nava, P. Coïsson, S. M. Radicella, "A new version of the NeQuick ionosphere electron density model", Journal of Atmospheric and Solar-Terrestrial Physics (2008), doi:10.1016/j.jastp.2008.01.015



Epstein function







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The model is represented by a sum of Epstein functions for the E, F1 and F2 layers:

$$N_{bot}(h) = N_E(h) + N_{F1}(h) + N_{F2}(h)$$

where

$$N_E(h) = \frac{4Nm^*E}{\left(1 + \exp\left(\frac{h - hmE}{BE}\xi(h)\right)\right)^2} \exp\left(\frac{h - hmE}{BE}\xi(h)\right)$$
$$N_{F1}(h) = \frac{4Nm^*F1}{\left(1 + \exp\left(\frac{h - hmF1}{B1}\xi(h)\right)\right)^2} \exp\left(\frac{h - hmF1}{B1}\xi(h)\right)$$
$$N_{F2}(h) = \frac{4NmF2}{\left(1 + \exp\left(\frac{h - hmF2}{B2}\right)\right)^2} \exp\left(\frac{h - hmF2}{B2}\right)$$



with

$$Nm^*E = NmE - N_{F1}(hmE) - N_{F2}(hmE)$$

 $Nm^*F1 = NmF1 - N_E (hmF1) - N_{F2} (hmF1)$

and

$$\xi(h) = \exp\left(\frac{10}{1+1|h-hmF2|}\right)$$

is a function that ensures a "fade out" of the E and F1 layers in the vicinity of the F2 layer peak in order to avoid secondary maxima around hmF2.



The model topside is represented by a semi-Epstein layer with a height-dependent thickness parameter H:

$$N(h) = \frac{4NmF2}{\left(1 + \exp\left(z\right)\right)^2} \exp\left(z\right)$$

with

$$z = \frac{h - hmF2}{H}$$
$$H = H_0 \left[1 + \frac{rg(h - hmF2)}{rH_0 + g(h - hmF2)} \right]$$



Peak heights

hmE = 120 $hmF1 = \frac{hmE + hmF2}{2}$ $hmF2 = \frac{1490MF}{M + \Delta M} - 176$

with

$$\Delta M = \begin{cases} 0.253/(foF2/foE) \\ -1.215) - 0.012, \\ -0.012 & \text{if } foE = 0, \end{cases}$$
$$MF = M \sqrt{\frac{0.0196M^2 + 1}{1.2967M^2 - 1}} \qquad M = M(3000)F2.$$

Thickness parameters

 $BE_{bot} = 5$

$$BE_{top} = \max\left(0.5\left(hmF1 - hmE\right), 7\right)$$

 $B1_{bot} = 0.5 \left(hmF1 - hmE \right)$

$$B1_{top} = 0.3 \left(hmF2 - hmF1 \right)$$

$$B2_{bot} = \frac{0.385NmF2}{(dN/dh)_{max}}$$
$$H = kB2_{bot} \left[1 + \frac{rg(h - hmF2)}{rkB2_{bot} + g(h - hmF2)} \right]$$





$$k = 3.22 - 0.0538 foF2 - 0.00664 hmF2 + 0.113 \frac{hmF2}{B2_{bot}} + 0.00257 R12 (M_{bot}) + 0.00257 R1$$



Critical frequencies and propagation factor

$$(foE)^{2} = \left(a_{e}\sqrt{F107}\right)^{2} \left(\cos\chi_{eff}\right)^{0.6} + 0.49$$

$$foF1 = \begin{cases} 1.4 \ foE & \text{if} & foE \ge 2\\ 0 & \text{if} & foE < 2 \end{cases}$$

 $0.85 \ 1.4 \ foE$ if $1.4 \ foE > 0.85 \ foF2$

foF2modeled in terms of ITU - R coefficientsM = M(3000)F2modeled in terms of ITU - R coefficients



Data ingestion/assimilation



- Empirical models like NeQuick have been developed as climatological models, able to reproduce the typical median condition of the ionosphere.
- For research purposes and practical applications, in order to pass from "climate" to "weather", there is a need to have models able to reproduce the current conditions of the ionosphere.
- Considering that there is an increasing availability of experimental data even in real time (ground and space-based GNSS, ionosondes), several assimilation schemes have been developed. They are of different complexity and rely on different kinds of data.



Data Assimilation

- "Data assimilation is an analysis technique in which the observed information is accumulated into the model state by taking advantage of consistency constraints with laws of time evolution and physical properties" (Bouttier, and Courtier, 1999).
- "Data assimilation is fundamentally a model specification and prediction technique that uses data to improve the fidelity of the model" (Bust and Mitchell, 2008).
- The ionospheric data assimilation schemes are based on different mathematical techniques; the most widely used are:
 - variational techniques (3D-VAR and 4D-VAR)
 - Kalman filters (extended Kalman filter (EKF), Ensemble Kalman filter (EnKF))



 Utah State University (USU) Global Assimilation of Ionospheric Measurements (GAIM) [Schunk et al., 2004] or the Jet Propulsion Laboratory (JPL)/University of Southern California (USC) Global Assimilative Ionospheric Model (GAIM) [Wang et al., 2004], or [Schunk et al., 2014], for example, are based on assimilation of data originating from different sources and imply the use of first principle models.

Midlatitude to Low-Latitude lonosphere Midlatitude to Low-Latitude lonosphere Midlatitude to Low-Latitude lonosphere with Drivers latitude to Low Latitude lonosphere-Plasmasphere with Drivers Midlatitude to Low Latitude lonosphere with Drivers High-Latitude lonosphere with Drivers Global Thermosphere Model-Data Assimilation
•

lonosphere	Electrodynamics	Thermosphere
Ground-based GPS-TEC	Ground magnetometers	Satellite UV emissions
Satellite-based GPS occultation	DMSP cross-track velocities	In situ neutral densities and winds
Ionosonde and digisonde	SuperDARN line-of-sight velocities	Satellite accelerometer and drag
In situ N _e	Iridium magnetometers	FPI winds
911 Å, 1356 Å, limb, disk (UV)	ACE interplanetary magnetic field, <i>Dst</i>	ISR neutral parameters
Solar UV, EUV	Solar UV, EUV	Solar UV, EUV



- The Electron Density Assimilative Model (EDAM) [Angling and Khattatov, 2006; Angling, M. J., and N. K. Jackson-Booth, 2011] provides a mean to assimilate ionospheric measurements into a background ionospheric model.
- Assimilated data are: ground-based and space-based GPS-derived TEC, ionosondes-derived parameters
 - Currently IRI is used as a background model (electron density only)
- Extended, localised Gauss Markov Kalman Filter
 - BLUE + time evolution of the differences between the measurements and the background ionosphere
 - Model variances are propagated
 - Covariance are estimated as required





- The Multi Instrument Data Analysis System (MIDAS) [Mitchell C. N. and Spencer P. S. 2003] is a tomographic approach where TEC data are inverted to evaluate the distribution and time evolution of electron concentration.
 - Orthonormal basis functions and SVD are used to solve the inverse problem.
- Review paper: Bust, G. S., and C. N. Mitchell (2008), "History, current state, and future directions of ionospheric imaging, Rev. Geophys., 46,RG1003, doi:10.1029/2006RG000212.



 TOMographic IONosphere model (TOMION), [Hernández-Pajares, M. et al., 1999] generates Global Ionospheric Maps (GIMs) of vertical TEC and includes an interpolation module using Kriging technique [Orús et al., 2005]. The ionosphere is represented by two or more layers of voxels and in each voxel the electron density is assumed to be constant. No background model is used.





 IRI Real Time Assimilative Model (IRTAM) [Galkin, I. A., et al. 2012], has been developed to assimilate Global Ionosphere Radio Observatory (GIRO) data (foF2, hmF2) in order to "update" the IRI electron density distribution, while preserving the IRI's typical ionospheric feature representations.

 The technique calculates the corrected coefficients for the spherical/diurnal expansion used by the CCIR-67/URSI-88 model to specify the global foF2 maps, and similarly the maps for all other IRI profile parameters.

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- A similar approach has been used by Brunini et al., [2013] in order to update the ITU-R database using radio occultation (COSMIC) electron density profiles.
- For this purpose the La Plata Ionospheric Model (LPIM) (after linearisation) is adjusted by Least Squares to every RO profile available for the time period of interest.



Global representation of the NmF2 estimated value within the 18–20 UT interval a) NmF2 for the 2007 September equinox c) NmF2 for the 2011 December solstice



- The Ionospheric Data AssimilationThree-Dimensional (IDA3D), [Bust et al., 2004] uses a three-dimensional variational data assimilation technique (3DVAR).
- It is capable of incorporating ground based and space based GPS-TEC measurements and electron density measurements from radars and satellites.
- The background specification is based upon empirical ionospheric models, but IDA3D is capable of using any global ionospheric specification as a background. IDA3D produces a spatial analysis of the electron density distribution at a specified time. A time series of these specifications can be created using past specifications to determine the background for the current analysis.



TRAIN – A realtime ionospheric DA model

- TRAIN is based on the Malvern Mathematical Data Assimilation Model (MMaDAM)
 - Generalised data assimilation system
 - Hybrid approach of ensemble Kalman filter (EnKF) and particle filter (PFs)
- Background model
 - Ensemble of NeQuick models
- Observation operator
 - Direct integration through each ensemble member
- Forward operator
 - Collapse ensemble to the mean
 - ICTP method for NeQuick forecast



Total Electron Content (TECU)

TRAIN electron density grid for 1200 UT, 4/6/2017 assimilating TEC data from ground based GPS receivers, a shipborne GPS receiver and a space based dual frequency altimeter.





Use of effective parameters

Pignalberi, A., Pezzopane, M., Rizzi, R. Galkin, I., "Effective Solar Indices for Ionospheric Modeling: A Review and a Proposal for a Real-Time Regional IRI"; Surv Geophys (2018) 39: 125. <u>https://doi.org/10.1007/s10712-017-9438-y</u>.



Effective parameters

- One of the first effective parameter that has been proposed is the "effective sunspot number" (SSNe). This parameter valid for a set of foF2 observations has been defined as the SSN value that, when used as input to the URSI foF2 model, gives a weighted zero-mean difference between the observed and the modeled foF2 values.
- T-index

http://www.sws.bom.gov.au/Educational/5/2/1 The T index can be best regarded as an "equivalent sunspot number" - the sunspot number which would best match the observations made by ionosondes.



Effective parameters

• IRI IG 12

http://gauss2.gge.unb.ca/grads/attila/papers/52am/ion52am.pdf An effective sunspot number (inferred IG12 index) is used which is defined as the product of the IG12 index and the scaling factor

 Klobuchar-Style Ionospheric Coefficients http://aiuws.unibe.ch/ionosphere/#cgim Klobuchar-style alpha and beta coefficients best fitting VTEC data

Effective parameters can be defined using diverse kind of data/ models



Basic ideas

NeQuick features relevant to implement adaptation techniques based on the use of effective parameters.

• For a given epoch & ray-path the model TEC is a monotonic function of the solar activity index, that can be regarded as an "effective ionization level" parameter.

✓ link, minimize TEC mismodeling → effective F10.7 (grid)
 ↓
 Use NeQuick to reconstruct the 3D electron density of the



vTEC map



grid points: lat.=-90°, 90° step 2.5° lon.=-180°, 180° step 5°

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Reconstructed foF2 map



grid points: lat.=-90°, 90° step 2.5° lon.=-180°, 180° step 5°

NeQuick2: validation results (example: HSA)



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STEC data ingestion; error statistics (000405)



foF2 data ingestion; error statistics, Sep 2011



Adapting NeQuick model to experimental slant TEC and foF2 data at a given location

(Use of slab thickness to constrain the NeQuick profile shape parameter)







Remarks

- The use of two effective parameters has been considered in order to use the ITUR coefficients to estimate foF2 and hmF2 in a region surrounding the ground station.
- In this way the peak parameter values can be estimated for a slant TEC computation.

Use JRO profiles to simulate the process of adapting NeQuick to GPS derived TEC and ionosonde peak parameters data.

TEC and peak parameters are known from the profile.

After model adaptation it is possible to compare profiles in order to evaluate the adaptation technique effectiveness.



Adaptation method validation



Jicamarca Radio Observatory (JRO) location

Adaptation method validation

Model: NeQuick

NeQuick adaptation to foF2 & hmF2

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NeQuick adaptation to foF2 & hmF2

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NeQuick adaptation to foF2, hmF2 (and Tautop)

Best Linear Unbiased Estimator (BLUE)*

x_t true model state (dimension *n*) **x**_b background model state (dimension *n*) **x**_a analysis model state (dimension *n*) **y** vector of observations (dimension *p*) *H* observation operator (dimension $p \ge n$) **B** covariance matrix of background errors $\varepsilon_{b} = (\mathbf{x}_{b} - \mathbf{x}_{t})$ (dimension $n \ge n$) **R** covariance matrix of observation errors $\varepsilon_{o} = (\mathbf{y} - \mathbf{H}[\mathbf{x}_{t}])$ (dimension $p \ge p$) **A** covariance matrix of analysis errors $\varepsilon_{a} = (\mathbf{x}_{a} - \mathbf{x}_{t})$ (dimension $n \ge n$)

*<u>https://www.ecmwf.int/sites/default/files/elibrary/2002/16928-data-assimilation-concepts-and-methods.pdf</u>

- The following hypotheses are assumed:
 - Linearized observation operator: the variations of the observation operator in the vicinity of the background state are linear:
 - ∀ x close enough to x_b, H(x) H(x_b) = H(x-x_b) where H is a linear operator
 - Non-trivial errors: **B** and **R** are positive definite matrices.
 - Unbiased errors: the expectation of the background and observation errors is zero:
 - $E[\mathbf{x}_{b}-\mathbf{x}_{t}] = E[\mathbf{y}-\mathbf{H}[\mathbf{x}_{t}]] = 0$

- The following hypotheses are assumed (cont.):
 - Uncorrelated errors: observation and background errors are mutually uncorrelated.
 - $E[(\mathbf{x}_{b}-\mathbf{x}_{t})(\mathbf{y}-\mathbf{H}[\mathbf{x}_{t}])^{T}] = 0$
 - Linear analysis: we look for an analysis defined by corrections to the background which depend linearly on background observation departures.
 - Optimal analysis: we look for an analysis state which is as close as possible to the true state in an r.m.s. sense (i.e. it is a minimum variance estimate).

The optimal least-square estimator (BLUE analysis) is defined by

 $\mathbf{x}_{a} = \mathbf{x}_{b} + \mathbf{K} (\mathbf{y} - \mathbf{H}[\mathbf{x}_{b}])$ $\mathbf{K} = \mathbf{B}\mathbf{H}^{\mathsf{T}}(\mathbf{H}\mathbf{B}\mathbf{H}^{\mathsf{T}} + \mathbf{R})^{-1}$ $\mathbf{A} = (\mathbf{I}\text{-}\mathbf{K}\mathbf{H})\mathbf{B}$

K is called *gain* of the analysis

The BLUE analysis is equivalently obtained as a solution to the variational optimization problem:

$$\mathbf{x}_{a} = \operatorname{Arg\,min} J$$

$$J(\mathbf{x}) = (\mathbf{x} - \mathbf{x}_{b})^{\mathrm{T}} \mathbf{B}^{-1} (\mathbf{x} - \mathbf{x}_{b}) + (\mathbf{y} - H[\mathbf{x}])^{\mathrm{T}} \mathbf{R}^{-1} (\mathbf{y} - H[\mathbf{x}])$$

$$= J_{b}(\mathbf{x}) + J_{o}(\mathbf{x})$$

where J is called the cost function of the analysis J_b is the background term; J_o is the observation term

If the background and and observation errors pdfs are Gaussian, then \mathbf{x}_a is also the maximum likelihood estimator of \mathbf{x}_t

In our case:

y ~ TEC
x_a ~ retrieved electron density
x_b ~ background electron density
H ~ "crossing lengths" in "voxels"

e.g. bckg_TEC_i = $\mathbf{H}_i \mathbf{x}_b = \sum_j H_{ij} x_{bj}$

 $\begin{array}{ll} R_{ij} = c_R \, \delta_{ij} \, y_i^2 & (measurements \ are \ independent) \\ B_{ij} = c_B \, x_{bi} \, x_{bj} \, Exp[-(z_{ij}/Lz)^2] \, Exp[-(\alpha_{ij}/L\alpha)^2] & (V \ \& \ H \ correl. \ are \ separable) \end{array}$

 z_{ij} ~ height difference between voxels i and j α_{ij} ~ angular (great circle) distance between voxels i and j

Lz ~ correl. distance in vert. direction (may depend on height) Lα ~ correl. (angular) distance in hor. direction (may depend location,...)

Data used

- For the assimilation
 - Ground-based GNSS-derived slant TEC data
 - Radio-Occultation-derived TEC data (calibrated TEC values along the LEO-to-GPS link below the LEO orbit)

- For the validation
 - Manually scaled ionosonde derived foF2 and hmF2
 - Ground-based GNSS-derived slant TEC data (not used for assimilation)

LISN: 3 days data (2011/03/11-12-13)

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Background 2011 03 11 UT: 15. -64.75°E

Assimilation effect

Background model (before the assimilation)

> Cross section 19:33UT; -64.75°E from -40°N to 10°N

Analysis (after the assimilation)

Ground-based & RO TEC DA

NeQuick used as a background with f10.7 input

Ne profiles at lonosonde location

2.0×10¹²

2.5×10¹²

3.0×10¹²

ionPhs C001.2011.070.23.31.G15 2010.2640 nc Peak Lon: -65.470° Peak UT: 23.470 Peak LT: 19.110 Peak dens: 1.70149700*10^12 m^-3 Peak freg: 11.710 MHz

Ground-based & RO TEC

1.5×10¹²

N(h) [m-3]

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TEC DA with NeQuick pre-adaptation

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Thank you for your attention

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